Geosci. Model Dev., 9, 111–124, 2016 www.geosci-model-dev.net/9/111/2016/ doi:10.5194/gmd-9-111-2016 © Author(s) 2016. CC Attribution 3.0 License.





Prediction of cloud condensation nuclei activity for organic compounds using functional group contribution methods

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Received: 31 July 2015 – Published in Geosci. Model Dev. Discuss.: 1 September 2015 Revised: 4 December 2015 – Accepted: 16 December 2015 – Published: 19 January 2016

Abstract. A wealth of recent laboratory and field experiments demonstrate that organic aerosol composition evolves with time in the atmosphere, leading to changes in the influence of the organic fraction to cloud condensation nuclei (CCN) spectra. There is a need for tools that can realistically represent the evolution of CCN activity to better predict indirect effects of organic aerosol on clouds and climate. This work describes a model to predict the CCN activity of organic compounds from functional group composition. Following previous methods in the literature, we test the ability of semi-empirical group contribution methods in Köhler theory to predict the effective hygroscopicity parameter, kappa. However, in our approach we also account for liquidliquid phase boundaries to simulate phase-limited activation behavior. Model evaluation against a selected database of published laboratory measurements demonstrates that kappa can be predicted within a factor of 2. Simulation of homologous series is used to identify the relative effectiveness of different functional groups in increasing the CCN activity of weakly functionalized organic compounds. Hydroxyl, carboxyl, aldehyde, hydroperoxide, carbonyl, and ether moieties promote CCN activity while methylene and nitrate moieties inhibit CCN activity. The model can be incorporated into scale-bridging test beds such as the Generator of Explicit Chemistry and Kinetics of Organics in the Atmosphere (GECKO-A) to evaluate the evolution of kappa for a complex mix of organic compounds and to develop suitable parameterizations of CCN evolution for larger-scale models.

1 Introduction

Organic compounds are an important contributor to the atmospheric submicron aerosol (Jimenez et al., 2009). The organic fraction is projected to increase in the future due to the confluence of a decreasing sulfate and nitrate burden and increases in the global secondary organic aerosol burden (Heald et al., 2008). An important unanswered question is how the organic influences the aerosol's ability to serve as cloud condensation nuclei (CCN), and in turn modulate climate via indirect effects of aerosols on clouds and precipitation (Andreae and Rosenfeld, 2008). Realistic prescribed variations in secondary organic aerosol hygroscopicity have demonstrable impacts on CCN number concentration (Mei et al., 2013) and can change the simulated global aerosol indirect forcing (AIF) by approx. one-sixth of the AIF simulated in a control case (Liu and Wang, 2010). To obtain a prognostic understanding of the contribution of the organic fraction to indirect aerosol forcing in future climates, models need improved schemes that map simulated organic aerosol composition to hygroscopicity and CCN activity.

Several organic aerosol types (e.g., freshly emitted diesel oil particles or first generation oxidation products of sesquiterpenes) consist of mostly hydrophobic hydrocarbon chains with few functional groups attached. Pure hydrocarbons with a carbon number less than C_{30} are expected to be semi-volatile and in the liquid phase. Over time the compounds evolve by functionalization, fragmentation, and oligomerization (Kroll and Seinfeld, 2008; Ziemann and Atkinson, 2012). As functional groups are added to the carbon chain, the products usually, but not always, become less volatile (Goldstein and Galbally, 2007), more dense (Kuwata et al., 2012), more viscous (Sastri and Rao, 1992), and more CCN active (Suda et al., 2014).

Laboratory (George and Abbatt, 2010; Poulain et al., 2010; Cappa et al., 2011; Massoli et al., 2010; Lambe et al., 2011; Duplissy et al., 2011; Kuwata et al., 2013; Rickards et al., 2013; Suda et al., 2014) and field studies (Jimenez et al., 2009; Chang et al., 2010; Mei et al., 2013) have demonstrated a robust link between the aerosol oxidation state and the ability of the organic fraction to promote hygroscopic water uptake and CCN activity. Proxies from mass spectrometry such as the fragmentation peak f_{44} or the atomic oxygento-carbon ratio are often used to model the increase in hygroscopicity. However, these correlations exhibit significant variability between studies and break down when applied at the compound level (Rickards et al., 2013; Suda et al., 2014).

Chemistry models are already capable of simulating the molecular identities of species present in the condensed phase during multi-day evolution of diluting air parcels (Lee-Taylor et al., 2015). Mapping this speciated aerosol composition to the aerosol hygroscopicity should ultimately permit quantification of changes in CCN number concentration (provided that the size distribution is also simulated) and associated effects on clouds and climate. Thermodynamic models should be able to predict CCN activity. Many thermodynamic models have made use of activity coefficients predicted by the universal functional group activity coefficient (UNIFAC) group contribution method (Fredenslund et al., 1975). Several investigators have compared UNIFAC predictions of organic aerosol water content to experimental data (Saxena and Hildemann, 1997; Ming and Russell, 2001; Peng et al., 2001; Choi and Chan, 2002; Mochida and Kawamura, 2004; Marcolli and Peter, 2005; Moore and Raymond, 2008). Some of these comparisons prompted proposed revisions of specific group interaction parameters, e.g., [OH] and [H₂O]. Several thermodynamic models that treat complex phase equilibria of multifunctional, multicomponent organic mixtures are based on UNIFAC activity coefficients (Ming and Russell, 2002; Raatikainen and Laaksonen, 2005; Topping et al., 2005; Amundson et al., 2007; Zuend et al., 2008; Compernolle et al., 2009). The development of these models has been driven by the need to enable predictions over a wide range of conditions and compositions, including the effect of liquid-liquid phase separation on gas-to-particle partitioning (Zuend and Seinfeld, 2012; Topping et al., 2013). The prediction of CCN activity of organic compounds has received less attention. Rissman et al. (2007) used the aerosol diameter-dependent equilibrium model (ADDEM; Topping et al., 2005) with an underlying UNIFAC core to predict the relationship between critical supersaturation and dry for several dicarboxylic acid aerosols. To our knowledge no study to date has systematically focused on the prediction of CCN activity from thermodynamic models.

Here we build on this body of work to predict the contribution of a compound with known chemical structure to the CCN activity of a particle of known size. The proposed model uses the UNIFAC equations (Fredenslund et al., 1975) with group interaction parameters form Hansen et al. (1991), Raatikainen and Laaksonen (2005), and Compernolle et al. (2009) to model activity coefficients and free energy of mixing. Liquid-liquid phase boundaries are determined using the area method of Eubank et al. (1992). Molecular volume is estimated from elemental composition and adjustments for functional group composition using the approach of Girolami (1994). The relationship between critical supersaturation and dry diameter is then predicted using Köhler theory (Seinfeld and Pandis, 2006). The basic model mechanics are similar to those employed in multicomponent phase equilibrium models (Ming and Russell, 2002; Raatikainen and Laaksonen, 2005; Topping et al., 2005; Amundson et al., 2007; Zuend et al., 2008) but limited in scope to binary compositions and with focus on accurately representing phase and water activity at conditions relevant at the point of CCN activation only. These predictions are validated by manually mapping chemical composition to UNI-FAC groupings and comparing modeled CCN activity against observations from a compiled library of recently published CCN data of mostly weakly oxidized hydrocarbons containing a mixture of alcohol, carbonyl, aldehyde, ether, carboxyl, nitrate, and hydroperoxide moieties. The model is used to predict how the addition of one or more functional groups to otherwise similar molecules promotes CCN activity. Envisioned application to multi-component aerosols and contrasts with more complete thermodynamic models are discussed.

2 Model description

2.1 Köhler theory

The saturation ratio over a curved droplet is given by the Köhler equation

$$S = a_{\rm w} \times \exp\left(\frac{4\sigma_{\rm s/a}(T)M_{\rm w}}{\rho_{\rm w}RTD}\right),\tag{1}$$

where a_w is the water activity, $\sigma_{s/a}$ is the surface tension of the solution/air interface, *T* is temperature, M_w is the molecular weight of water, ρ_w is the density of pure water, *R* is the universal gas constant, and *D* is the wet drop diameter. Water activity depends on the water content and the amounts and identities of solutes in the nucleus. The principle water content variable used in this work is the mole fraction

$$x_{\rm w} = \frac{n_{\rm w}}{n_{\rm w} + \sum_i n_{\rm s,i}},\tag{2}$$

where x_w is the mole fraction of water, n_w and $n_{s,i}$ are the number of moles of water and solutes, and *i* is the number of dry components. The wet drop diameter can be calculated from x_w if the dry diameter, D_d , is specified and it is assumed

that the particle is spherical and that the volume of water and solute are additive:

$$D = \langle (x_{\rm w} - 1)^{-1} \left(x_{\rm w} - x_{\rm w} \sum_{i} (\varepsilon_i v_{\rm w} v_{{\rm s},i}^{-1}) - 1 \right) D_{\rm d}^3 \rangle^{1/3}.$$
 (3)

In Eq. (3) v_w and $v_{s,i}$ are the molar volume of the water and solutes and ε_i are the volume fractions in the dry particle. Equation (3) is obtained by rearranging Eq. (7) in Petters et al. (2009a). The critical supersaturation required for an aqueous solution droplet to activate into a cloud droplet is found by combining Eqs. (1) and (3) and finding the x_w (or *D*) that maximizes s_c

$$s_{c} = \{\max[a_{w} \\ \times \exp\left(\frac{4\sigma_{s/a}(T)M_{w}}{\rho_{w}RT\langle (x_{w}-1)^{-1}\left(x_{w}-x_{w}\sum_{i}(\varepsilon_{i}v_{w}v_{s,i}^{-1})-1\right)D_{d}^{3}\rangle^{1/3}}\right)\right]\right\} \times 100\%$$
(4)
$$x_{w} \in [0, 1],$$

where s_c is the critical supersaturation in %. The variables that control s_c are v_s , a_w , and $\sigma_{s/a}$. In this work it is assumed that surface tension is that of pure water. Discussion on this and other assumptions are provided at the end of this section. First the prediction of v_s and a_w for organic compounds with known chemical structure is described.

2.2 Molar volume

Molar volume is calculated from the molecular formula using the method of Girolami (1994). Each element is assigned a relative volume based on its location in the periodic table. The elemental volumes are summed and scaled by a constant factor to compute v_s . If the oxygen is bound in the form of alcohol [OH] or carboxyl [C(=O)OH] moieties, the actual v_s is smaller due to intramolecular bonding. Therefore, v_s is decreased by 10% for each [OH] or [C(=O)OH] group but by no more than 30% of the molar volume derived from the elemental composition. Girolami (1994) tested this method for 166 liquids and reports agreement with observations $v_s \sim \pm 10$ %. Barley et al. (2013) reviewed the performance of various methods for predicting molar volume using a test set of 56 multifunctional organic compounds and report similar scatter.

2.3 Water activity

Water activity is related to the mole fraction via

$$a_{\rm w} = \gamma_{\rm w} x_{\rm w},\tag{5}$$

where γ_w is the activity coefficient of water. Activity coefficients are estimated using the semi-empirical group contribution method UNIFAC (Fredenslund et al., 1975). The UNIFAC model describes a liquid solution that consists of *i* components. Each component is divided into *k* groups. The activity coefficient of component *i* in solution (γ_i) has contributions from combinatorial (γ^C) and residual parts (γ^R)

$$\ln \gamma_i = \ln \gamma_i^{\rm C} + \ln \gamma_i^{\rm R}.$$
 (6)

The combinatorial part is computed via

$$\ln \gamma_i^{\rm C} = \ln \frac{\Phi_i}{x_i} + \frac{z}{2} q_i \ln \frac{\theta_i}{\Phi_i} + l_i - \frac{\Phi_i}{x_i} \sum_j x_j l_j, \tag{7a}$$

$$l_i = \frac{z}{2}(r_i - q_i) - (r_i - 1); \quad z = 10,$$
(7b)

$$\theta_i = \frac{q_i x_i}{\sum_j q_j x_j}; \quad \Phi_i = \frac{r_i x_i}{\sum_j r_j x_j}, \tag{7c}$$

$$r_i = \sum_k v_k^{(i)} R_k; \quad q_i = \sum_k v_k^{(i)} Q_k.$$
 (7d)

In Eqs. (7), x_i is the mole fraction of component i, θ_i , and Φ_i are the average surface and segment fraction, z is the lattice coordination number, $v_k^{(i)}$ is the number of groups of type k in component i, R_k , and Q_k are the group volume and surface area parameters derived from Bondi (1964), and r_i and q_i are the normalized van der Waals volume and surface area. The summation i or j is over all components in the mixture, including component i.

The residual part is computed via

$$\ln \gamma_i^{\mathbf{R}} = \sum_k v_k^{(i)} \left[\ln \Gamma_k - \ln \Gamma_k^{(i)} \right], \tag{8a}$$

$$\ln \Gamma_{k} = Q_{k} - \left[1 - \ln\left(\sum_{m} \Theta_{m} \Psi_{mk}\right) - \sum_{m} \frac{\Theta_{m} \Psi_{km}}{\sum_{n} \Theta_{n} \Psi_{nm}}\right],$$
(8b)

$$\Theta_m = \frac{Q_m X_m}{\sum_n Q_n X_n},\tag{8c}$$

$$X_m = \frac{\sum_i v_m^{(i)} x_i}{\sum_i \sum_k v_k^{(i)} x_i},$$
(8d)

$$\Psi_{mn} = \exp\left(-\frac{a_{mn}}{T}\right). \tag{8e}$$

In Eqs. (8), a_{mn} are empirically determined parameters, Ψ_{mn} is the group interaction parameter of group *m* with *n*, X_m is the mole fraction of group *m* in the mixture, Θ_m is the area fraction of group *m*, Γ_k is the group residual activity coefficient, and $\Gamma_k^{(i)}$ is the residual activity coefficient of group *k* in a reference solution containing only molecules of type *i*. Equations (8) are also used to compute $\Gamma_k^{(i)}$. The summation *n* or *m* is over all different groups in the mixture, and the summation *k* is over all groups in component *i*.

Groups within UNIFAC are represented as main groups and subgroups. The main groups evaluated in this work

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are alkane [CH_n], alcohol [OH], water [H₂O], carbonyl [CH_nC(=O)], aldehyde [HC(=O)], ether [CH_n(O)], carboxyl [C(=O)OH], nitrate [CH_nONO₂], and hydroperoxide [CH_n(OOH)]. Interaction parameters a_{mn} between the main groups that are used in this work are tabulated in Table S1 in the Supplement. Some of the main groups have several subgroups, with each subgroup having unique volume and surface area parameters R_k and Q_k . These are summarized in Table S2.

2.4 Phase equilibrium

For some x_w liquid–liquid phase separation can occur. The normalized Gibbs free energy of the mixture, defined as the actual Gibbs free energy divided by the thermal energy, is needed to compute the number of thermodynamically stable phases in the system. For a binary system consisting of water (w) and a single solute (s), Gibbs energy is calculated from the activity coefficients via standard thermodynamic relationships (Prausnitz et al., 1999; Petters et al., 2009a)

$$\Delta g^{\rm mix} = \Delta g^{\rm ideal} + \Delta g^{\rm excess},\tag{9a}$$

$$\Delta g^{\text{ideal}} = x_{\text{w}} \ln x_{\text{w}} + (1 - x_{\text{w}}) \ln x_{\text{s}}, \qquad (9b)$$

$$\Delta g^{\text{excess}} = x_{\text{w}} \ln \gamma_{\text{w}} + (1 - x_{\text{w}}) \ln \gamma_{\text{s}}, \qquad (9c)$$

where Δg^{mix} is the normalized change in Gibbs free energy of the mixture, Δg^{ideal} is the change in ideal Gibbs free energy of the mixture (Raoult's law), and Δg^{excess} is the excess Gibbs free energy of mixing quantifying the deviation from Raoult's law. In highly non-ideal solutions liquid–liquid phase separation may occur. Two compositions x_a and x_b define the water mole fraction of the two co-existing phases. Computationally, x_a and x_b can be obtained from Δg^{mix} using the area method (Eubank et al., 1992). Briefly, the state space is evaluated by computing the following area for all possible combinations x_I and x_{II}

$$A(x_{\mathrm{I}}, x_{\mathrm{II}}) = \left| \left[\Delta g^{\mathrm{mix}}(x_{\mathrm{II}}) + \Delta g^{\mathrm{mix}}(x_{\mathrm{I}}) \right] \left[\frac{x_{\mathrm{II}} - x_{\mathrm{I}}}{2} \right] \right| - \left| \int_{x_{\mathrm{I}}}^{x_{\mathrm{II}}} \Delta g^{\mathrm{mix}}(x) \mathrm{d}x \right|.$$
(10a)

Phase boundaries x_a and x_b exist if condition

$$A(x_a, x_b) = \max A(x_{\mathrm{I}}, x_{\mathrm{II}}); \quad A > 0$$
(10b)

is satisfied. If multiple phases coexist in phase equilibrium, the Gibbs–Duhem relationship dictates that the chemical potential of each component is equal in all phases. Therefore the water activity inside the miscibility gap is constant and the values entering Eq. (4) are subject to the constraint

$$a_{w} = \begin{cases} a_{w}(x_{a}) = a_{w}(x_{b}) & \text{for } x_{a} \le x_{w} \le x_{b} \\ \gamma_{w}x_{w} & \text{else.} \end{cases}$$
(11)

We note that Eubank et al. (1992) algorithm can be extended to *n* components. Other numerically efficient approaches to find phase equilibrium, including those of *n* component mixtures, are available in the literature (e.g., Amundson et al., 2005, 2007; Zuend et al., 2010). Comparison for phase boundaries (x_a , x_b) calculated using standard UNIFAC parameters and the Eubank method used in this model, and standard UNIFAC parameter and the algorithm in the UHAERO model (Amundson et al., 2007) are in good agreement and summarized in the Supplement.

2.5 Model implementation

The model was implemented to run on a personal computer using the commercial MATLAB environment (Math-Works, Inc.). Alternatively, the code runs under the Octave environment, which is available as free software under the GNU General Public License. Correct implementation of the UNIFAC model was confirmed by comparing results from test mixtures against output from existing implementations, which is further described in the Supplement. A compound is defined by specifying a count of subgroups comprising the molecule. Equations (6)-(8) are solved to find $\gamma_{\rm w}$ for *n* linearly spaced values within the domain $x_{\rm w} \in$ [0.0001, 0.9999]. Resulting γ_w are parsed through Eqs. (9)– (11) to find the number of stable phases and to define a_w over the entire domain. These a_w are interpolated onto a higher resolution linearly gridded domain (m points) to improve the accuracy of the computation of s_c using Eq. (4). Values for *n* and *m* are selected to balance computational speed and solution accuracy. Equations (6)-(8) have linear time complexity. Equations (9)–(11) have quadratic time complexity. Thus, the two algorithms have an order of O(n) and $O(n^2)$, respectively. For n > 200, the overall model time complexity is $O(n^2)$. For $n > \sim 800$ and m = 10000, the resolution is sufficiently high so that the computed s_c becomes independent of the choice of n. All computations in this work were carried out for n = 1000 and m = 10000. Total model execution times for a single compound on an Intel(R) Core(TM) i7-2600 3.4 GHz microprocessor using a single core were 39 s with MATLAB version R2013a (8.1.0.604) 64 bit and 282 s with GNU Octave version 3.8.1 configured for 64 bit.

2.6 Hygroscopicity parameter

Equation (4) is solved to find s_c for a specified dry diameter, fixed T = 298.15 K and $\sigma_{s/a} = 0.072$ J m⁻². The result is expressed in terms of the hygroscopicity parameter κ (Petters and Kreidenweis, 2007) that is defined via

$$s_{\rm c} = \left\{ \max\left[\frac{D^3 - D_{\rm d}^3}{D^3 - D_{\rm d}^3(1 - \kappa)} \exp\left(\frac{4\sigma_{\rm s/a}}{\rho_{\rm w}RTD}\right) \right] - 1 \right\} \times 100\% \quad (12)$$
$$D \in [D_{\rm d}, \infty].$$

The hygroscopicity parameter is obtained by iteratively seeking the κ value that satisfies Eq. (12) for a given D_d , s_c pair.

Kappa values obtained by fitting a D_d , s_c pair to Eq. (12) with the assumed temperature and surface tension conceptually correspond to "apparent hygroscopicity at standard state" (Christensen and Petters, 2012). All values in this work are apparent κ 's. For simplicity these are denoted as κ without further qualification. Observations against which the model is evaluated are summarized in the Supplement and will be discussed further in Sect. 3.

2.7 Model assumptions and limitations

The model approach presented here is limited to liquid organic compounds. This assumption is implied in both molar volume and UNIFAC activity coefficient calculations. Comparison with observational CCN data where the reference phase state may be crystalline should be interpreted with caution. For example, CCN experiments performed with crystalline dicarboxylic acids demonstrate that for some compounds deliquescence, i.e., a solubility-controlled phase transition, must precede droplet activation (Petters and Kreidenweis, 2008). The UNIFAC approach is unable to accurately predict the solubility of these compounds if they existed in their crystalline solid state. If, however, the compound is in metastable aqueous solution, the UNIFAC prediction is expected to be valid to within the general accuracy of the specific model implementation. Under atmospheric conditions where the organic compounds are embedded in a matrix comprising a multitude of organic compounds, a liquid or amorphous solid is the prevailing stable phase (Marcolli et al., 2004). Furthermore, since metastable states with hygroscopically bound water appear to dominate in the atmosphere (Rood et al., 1989; Nguyen et al., 2014) the liquid assumption may not be a serious limitation. Nonetheless, it is unclear whether the assumption of a liquid-like reference state is a serious limitation if the organic particles are highly viscous (Vaden et al., 2011; Shiraiwa et al., 2011; Zobrist et al., 2011; Renbaum-Wolff et al., 2013).

Other limitations of the UNIFAC method are the problems of accounting for group proximity effects and the inability to distinguish between isomers. Proximity effects occur when polar groups are separated by less than three to four carbon atoms (Topping et al., 2005). Since only the number of groups of type *i* are specified, all isomers are modeled to have identical κ values. Although experiments show that the location of the functional group has a small and systematic effect on the observed κ (Suda et al., 2014), those effects are relatively small and beyond the resolution of the model presented here.

The application of Eq. (4) assumes that the surface tension is that of pure water. Many organic compounds found in ambient organic aerosol lower the surface tension at the solution–air interface (Tuckermann and Cammenga, 2004; Tuckerman, 2007). However, several studies have demonstrated via experiment and theory that surfactant partitioning between the bulk solution and the Gibbs surface phase greatly diminishes the effect one would predict by applying macroscopic surface tensions in Köhler theory (Li et al., 1998; Rood and Williams, 2001; Sorjamaa et al., 2004; Prisle et al., 2011). Neglecting to account for reduced surface tension and using water activity to estimate CCN activity results in an underestimate of κ by ~ 30 % for the strong surfactant sodium dodecyl sulfate (Petters and Kreidenweis, 2013). We note that estimates of surface tension reduction for pure organic liquids can be obtained from critical pressure and boiling point (Sastri and Rao, 1995) and the Sprow and Prausnitz (1966) expression coupled with UNIFAC activity coefficients (Topping et al., 2005; Rafati et al., 2011). Combined with predictions of critical properties from functional group data (Joback and Reid, 1987), predicted binary surface tensions could be obtained for each compound. Including surfactant partitioning in Eq. (4) is possible using the expressions in Petters and Kreidernweis (2013) or similar approaches (Sorjamaa et al., 2004; Raatikainen and Laaksonen, 2011). Thorough validation against experimental data, including measurements of surface tension and CCN activity, is needed before this approach should be adopted.

2.8 Relationship to other thermodynamic models and application to multicomponent systems

The basic model functionality described here can also be obtained by appropriately initializing other multicomponent equilibrium models (Ming and Russell, 2002; Raatikainen and Laaksonen, 2005; Topping et al., 2005; Clegg and Seinfeld, 2006; Amundson et al., 2007; Zuend et al., 2008) with a set of binary water/organic solutions, parsing the output through a phase equilibrium module (if not included in the thermodynamic model itself) and the Köhler model. The predicted CCN activity mostly depends on the underlying set of group interaction parameters. The output should match with the solution presented here if the same interaction parameter matrix is used. The main conceptual distinction between the approach proposed here and the approach employed by the more complex multicomponent models is our focus on predictions for binary organic/water solutions and limitation of the scope to a narrow range of water activities relevant to CCN activation only. Accurate representation of hygroscopic growth at $a_w < \sim 0.99$ is not required and would be of secondary concern when tuning interaction parameters.

We envision that the proposed specialized model approach can be used to categorize individual compounds into three miscibility regimes, analogous to the solubility regimes defined in Petters and Kreidenweis (2008). Regime I: the compound is CCN inactive and can be effectively modeled as $\kappa = 0$. Regime II: the compound is CCN active without any additional phase constraints. In turn κ is mostly determined by molar volume and slightly modulated by activity coefficients. Regime III: the compounds' CCN activity is limited due to miscibility constraints. In turn κ is highly sensitive to overall water content and can either have $\kappa \sim 0$ or express κ according to its molar volume. Once pure component κ 's are predicted and stored in a database, the overall organic aerosol (OA) κ in mixed particles can be calculated quickly using the volume-weighted mixing rule (Petters and Kreidenweis, 2007). This compound-by-compound treatment of multicomponent mixtures assumes that solute-solute interactions are negligible. Salting-in and salting-out of solution effects are not captured. Effective κ values for compounds falling into the limited miscibility regime may be misrepresented in this treatment. Whether such effects are important will depend on the fraction of compounds in a mixture that fall into the limited miscibility regime and whether the proposed approach of intermediate complexity - modeling binary solutions coupled with a linear mixing rule – ultimately proves sufficiently accurate to model the evolution of ambient OA. In the following we use experimental data to demonstrate that the outlined UNIFAC model is suitable to categorize compounds into these three regimes.

3 Results and discussion

Experimental data for validation were compiled from the literature. A detailed summary of the compound names, chemical structures, physicochemical properties, CCN observations, and observed κ_{app} 's is provided in the Supplement (Tables S3-S7). This set features compounds with mostly linear carbon backbones C_4 to C_{18} and O: C ratio between 0.1 and 1. The data are grouped into model compounds for primary organic aerosol (POA; Table S3), functionalized hydroperoxy ethers (Table S4), hydroxy nitrates (Table S5), carboxylic acids (Table S6), and carbohydrates (Table S7). Compounds included in Table S3 are long-chain molecules that have hydrophobic tails (>14 methylene groups) and a single terminal carboxyl or hydroxyl group. Representative example compounds are oleic acid or cetyl alcohol. Compounds in Table S4 are C_{14} functionalized hydroperoxy ethers that have 10-12 methylene groups, at least one hydroperoxide and ether group, and a second carbonyl, hydroperxide, or carboxyl group. Compounds in Table S5 are functionalized hydroxy nitrates featuring C10 to C15 carbon backbones with 1-3 hydroxyl and 1-4 nitrate groups. Compounds in Table S6 are C₄-C₁₀ carboxylic acids that have 1-2 carboxyl and up to one carbonyl group attached to the carbon backbone. Finally, compounds in Table S7 are C₄–C₁₈ carbohydrates that have hydroxyl groups approximately equal to the number of carbon atoms. Data in Table S3 are taken from Raymond and Pandis (2002) and Shilling et al. (2007). Data in Tables S4 and S5 are taken from the Supplement of Suda et al. (2014). Data in Tables S6 and S7 are from various sources and are summarized in the Supplement of Petters et al. (2009b), which was updated with new compounds from Christensen and Petters (2012), and data were re-screened for quality. The compounds were selected to provide systematic variation in the number and type of functional groups with otherwise similar structure, i.e., linear or weakly branched alkane backbone with variable carbon chain length.

To illustrate model initialization and model output, two example compounds from the Supplement, C12 dihydroxy nitrate and C₁₃ trihydroxy nitrate, are presented in Table 1. For some of the compounds density and solubility data are available and those data are included in the Supplement. Table 1 shows how the molecular structure is decomposed into the subgroups understood by the UNIFAC and Girolami (1994) model framework. Detailed model output for the two example compounds is illustrated in Fig. 1. The predicted mole fraction dependence of Δg_{mix} suggests that the C₁₃ trihydroxy nitrate is miscible with water in all proportions while the C₁₂ dihydroxy nitrate is not. The dashed black line connecting x_a and x_b encloses the maximum positive area with the Δg_{mix} line and defines the two-phase region. Water activity derived from Δg^{mix} is graphed in the middle panel. It shows that the miscibility gap for the C₁₂ dihydroxy nitrate occurs at water activity close to unity. Phase gaps at water activity near unity may result in miscibility-controlled cloud droplet activation (Petters et al., 2006), which is analogous to solubility-/deliquescence-limited cloud droplet activation (Shulman et al., 1996; Hori et al., 2003; Bilde and Svenningsson et al., 2004; Kreidenweis et al., 2006; Petters and Kreidenweis, 2008). Köhler curves in the right panel demonstrate miscibility-limited activation behavior. For the C13 trihydroxy nitrate, the Köhler curve is smooth and exhibits a single maximum corresponding to the model critical supersaturation. For the C12 dihydroxy nitrate two maxima appear. The first maximum corresponds to the point of incipient phase separation x_a . The height of the miscibility barrier depends on the dry diameter. For large dry particles where the Kelvin term does not play a significant role, the supersaturation of point x_a is reduced and the second classical Köhler maximum will control droplet activation. Similar complex Köhler curves have been reported previously (e.g., Bilde and Svenningsson, 2004; Petters and Kreidenweis, 2008). Experiments with pure crystalline sparingly soluble organic compounds have demonstrated convincingly that the larger maximum indeed controls cloud droplet activation for solubility-limited cases (Hori et al., 2003; Bilde and Svenningsson, 2004; Hings et al., 2008). The s_c vs. D_d relationship for phase-controlled activation does not result in κ_{app} that is independent with respect to D_d (Petters and Kreidenweis, 2008). Therefore, for compounds having $\kappa < \sim 0.06$ where phase separation might play a role, the observed s_c , $D_{\rm d}$ pair is included in the data tables (Tables 1, S3–S7) and κ values are computed from the observation and the model (Eq. 12) at the same D_d . Note that the D_d -dependent κ only plays a role in a narrow range of miscibilities. Sufficiently soluble and truly insoluble substances are not affected. In summary, Table 1 and Fig. 1 demonstrate model input, illustrate model mechanics, and identify model outputs.

How well do data-derived and model-derived κ_{app} compare? For numerical comparison both κ 's are included in Ta-

Table 1. Properties for two example chemical compounds. UNIFAC representation indicated the number and type of subgroups to represent the chemical structure: MW denotes molecular weight $(g \text{ mol}^{-1})$ and v_s denotes the model predicted molar volume $(\text{cm}^3 \text{ mol}^{-1})$. CCN reflects the observed supersaturation and dry diameter data pair obtained from the source (Suda et al., 2014) from which observed κ was determined.

Name	Formula	Structure	UNIFAC representation		MW	Ve	Observed CCN	Apparent κ	
			no.	subgroup		. 8	s _c (%)	observed	model
			101	subBroup			D _d (nm)	00501104	
C ₁₂ dihydroxy nitrate	C ₁₂ H ₂₅ O ₅ N	O2NO OH	2 8	CH ₃ CH ₂	263.3	263.3	0.3	0.018	0.008
			1 1 2	C CH(ONO ₂) OH			222		
C ₁₃ trihydroxy nitrate	C ₁₃ H ₂₇ O ₆ N	OH OH	2 8 1	CH ₃ CH ₂ CH	293.4	257.7	0.3	0.1	0.07
			1 1 3	C CH(ONO ₂) OH			111		



Figure 1. Modeled Δg_{mix} (left), water activity (middle), and Köhler curves (right) for C₁₂ dihydroxy nitrate and C₁₃ trihydroxy nitrate (see Table 1). Open circles denote the mole fractions x_a and x_b that correspond to the envelope of compositions where liquid–liquid phase separation is predicted for the C₁₂ dihydroxy nitrate.

bles S3–S7. A graphical illustration of these is presented in Fig. 2. To improve clarity, compounds with predicted and modeled $\kappa < 0.001$ are clustered in the lower left corner. Such low κ 's correspond to compounds that are effectively CCN inactive. The range between $\kappa = 10^{-3}$ and 10^{-5} spans a narrow range in the $s_c - D_d - \kappa$ state space that characterizes CCN activity (cf. Fig. 1 in Petters and Kreidenweis, 2007). Resolving these differences is not particularly meaningful for organic dominated particles that typically have $D_d < 300$ nm. Furthermore, the κ of an internally mixed particle is approximately the weighted volume fraction in the mixture. For $\kappa < 10^{-3}$ the contribution to a mixed particle's κ is insensitive to the exact value. Finally, although state-of-the-science size-resolved CCN measurements can resolve differences in $\kappa < 10^{-3}$, compound impurities can interfere. A 1 % impurity having κ similar to ammonium sulfate would contribute ~ 0.06 to a measured particle κ . In addition, solvent residuals (Huff Hartz et al., 2006; Shilling et al., 2007; Rissman et al., 2007) and control over the dry particle phase state (Raymond and Pandis, 2002; Hori et al., 2003; Broekhuizen et al., 2004; Bilde and Svenningson, 2004) can disproportionally bias the characterization of low κ 's. Combined these points justify the definition of $\kappa < 0.001$ as effectively CCN inactive. Compounds in the CCN inactive corner include all compounds from Table S3, the C₁₄ and C₁₅ hydroxnitrate, and the C₁₄ trinitrate. These compounds all have 11 or more methylene groups and O : C ratios between 0.11 and 0.65. CCN activity of these compounds is satisfactorily predicted by the model.

Nine compounds are predicted to be CCN inactive but have measurements indicating $0.001 > \kappa_{obs} > \sim 0.03$. These are graphed below the dashed line and include C₁₄ di- and tetra-nitrate, C₁₃ hydroxy nitrate, C₁₄ and C₁₅ dihydroxy nitrate, the remaining hydroperoxide ethers from Table S4, and cis-pinonic acid. The observed C₁₄ di- and tetra-nitrate are



Figure 2. Model predicted vs. experimentally determined κ values. Values $\kappa < 0.001$ are classified as CCN inactive and are clustered in the lower left corner of the graph. Colors are used to delineate the grouped source data in the Supplement. Selected structures from the Supplement are included in the graph. C_x HN, C_x DHN, and C_x THN denote hydroxy nitrate, dihydroxy nitrate, and trihydroxy nitrate and x denotes the total number of carbon atoms. C_{14} DiN, C_{14} TriN, C_{14} TetraN denote the C_{14} dinitrate, trintrate, and tetranitrate, respectively. Points below the dashed line corresponds to compounds with predicted $\kappa < 0.001$ and observed $\kappa > 0.001$. Typical range of observed κ CCN for peroxides is indicated by the horizontal bar.

barely larger than the cutoff for CCN inactive. Variation of κ between the C₁₄ di-, tri- and tetra-nitrate (cf. Fig. 2 in Suda et al., 2014) implies that the trinitrate has lower κ than the di- and tetra-nitrate, which suggests that some random variability in the data is superimposed on the trend. Similarly, the observations show that the C₁₄ and C₁₅ dihydroxy nitrate are slightly more CCN active than the C₁₃ dihydroxy nitrate. Although this is possible such behavior is not plausible due to the well-established hydrophobic nature of the added CH_x groups. One possible explanation for the discrepancies

is the sensitivity of observed κ 's to trace contamination. Each of the compounds was purified via high-performance liquid chromatography (HPLC; Suda et al., 2014) but degree of purification likely varied between compounds. Furthermore, experimental uncertainty for the HPLC-CCN method used is slightly larger than for standard methods since it requires application of fast-flow scans. Finally, the data are from a single set of experiments. More data are needed before attributing the mismatch to either model or measurement error.

Another notable outlier is adipic acid. Here, the observed $\kappa < 0.01$ corresponds to the solubility-limited value that is referenced against its solid crystalline phase state. In contrast, the predicted value $\kappa = 0.14$ is in good agreement with the molar volume prediction ($\kappa = 0.17$; cf. Fig. 4 in Christensen and Petters, 2012) and observed κ that adipic acid particles express when solubility limitations are removed (cf. Fig. 1 in Hings et al., 2008). This scenario was selected to illustrate the inability of the UNIFAC model to treat solid phases. It therefore cannot capture deliquescence and deliquescence-/solubility-limited activation. In atmospheric OA multiple organic compounds likely form an amorphous supercooled melt (Marcolli et al., 2004) and metastable aqueous solutions are ubiquitous (Rood et al., 1989). Thus the metastable prediction would be valid to account for adipic acid in the context of atmospheric OA.

A series of carboxylic acids and carbohydrates cluster near the 1:1 line at $\kappa > \sim 0.06$. These compounds are generally highly functionalized having at least two carboxyl, hydroxyl, or carbonyl groups for every four carbon atoms. The O: C ratio always exceeds 0.5 and is close to 1 for many of the compounds. For the predictions, activity coefficients approach unity, compounds are miscible in water in all proportions, and model κ 's closely track the prediction based on estimated molar volume. Overall comparison of predicted vs. observed κ is approximately within a factor of 2 and this range is similar to predictions that are based on actual molar volume (cf. Fig. 2 in Petters et al., 2009b).

The series of hydroxy nitrates, dihydroxy nitrates, and trihydroxy nitrates for different carbon chain lengths also clusters near the 1:1 line. The spread is within approximately a factor of 2 and similar to that of the carboxylic acids and carbohydrates. These compounds span the entire range from $\kappa < 0.001$ to $\kappa \sim 0.1$ and have as few as two hydroxyl and one nitrate group per 13 carbon atoms (C13 dihydroxy nitrate). The model appears to accurately predict the influence of the methylene and hydroxyl groups on the transition from immiscible and CCN inactive to sufficiently miscible and CCN active according to the molar volume of the compound. For the C_{11} , C_{12} , and C_{13} dihydroxy nitrates, the predicted miscibility-limited activation demonstrated in Fig. 1 seems to adequately explain the transition. The accurate model prediction of this sensitive transition regime is encouraging, especially since no adjustment was made to the a_{mn} group interaction parameters for [OH], $[CH_x]$, and $[H_2O]$ groups.

In summary, Fig. 2 demonstrates four capabilities of the model. First, the model has good skill in correctly classifying effectively CCN inactive compounds ($\kappa < 0.001$). Second, the model captures the molar volume-dependent activation of highly functionalized compounds (low molecular weight dicarboxylic acids and polysaccharides). Scatter between predicted and observed κ is approximately within a factor of 2 and considered acceptable taking into account the considerable diversity in the underlying CCN data. We note that uncertainties in molar volume estimation of $v_s \sim \pm 10\%$

stemming from the Girolami (1994) method correspond to $\pm 10\%$ error in predicted κ for these compounds, which is significantly less than the observed scatter in the data (Petters et al., 2009b). Third, the model predicts that miscibility limitations are the cause for poor CCN activity of weakly functionalized hydrocarbons, and the phase separation information can be used to quantitatively predict the transition between sufficiently miscible and effectively immiscible species. Finally, the model seems to accurately capture the main functional group dependencies observed previously (Suda et al., 2014): a strong promoting effect of hydroxyl, a weak promoting effect for hydroperoxides, a negligible or inhibiting effect of nitrate, and inhibiting effect of methylene groups on CCN activity. How, then, can one quantify the model sensitivity of κ to the addition of functional groups to otherwise similar molecules?

Simulation of homologous series can be used to derive these sensitivities. Figure 3 shows modeled κ 's for a series of functionalized *n*-alkanes. The gradual decreasing trend of κ with increasing carbon number is due to the increase in molar volume. A steep decline is observed when a critical carbon number is exceeded. Beyond this point the additional methylene groups reduce the miscibility with water and render the compound effectively CCN inactive. For example, CCN activity for a C_{16} trihydroxy alkane is controlled mostly by molar volume while C_{18} trihydroxy alkane is effectively CCN inactive. The critical carbon number is C_7 , C_{12} , C_{16} , C_{20} , and C_{24} for the mono-, di-, tri-, tetra-, and penta-hydroxy alkanes, respectively. Starting with an n-alkane, the most dramatic effect of adding functional groups is to render the molecule miscible with water. Contrasting the critical carbon number for different homologous series can be used as a measure of a particular groups' ability to transform the molecule such that it is sufficiently miscible in water and can express its molar volume κ . The hydroxy alkane series shows that approximately one hydroxyl group is needed to compensate for the addition of four methylene groups (i.e., to maintain miscibility at the composition of the critical carbon number), expressed as a ratio, $\Delta[CH_n]/\Delta[OH] \sim -4/1$. Similar ratios for the other groups are derived from the shifts in the dihydroxy alkane series upon further functionalization: Δ [CH_n]/ Δ [C(=O)OH] ~ -5/2, Δ [CH_n]/ Δ [CH_nC(=O)] $\sim -2/3$. $\Delta[CH_n] / \Delta[HC(=O)]$ $\sim -4/2$, $\Delta[CH_n] / \Delta[CH_n(O)] \sim -2/4, \quad \Delta[CH_n] / \Delta[CH_n(OOH)]$ $\sim -2/2$. $\Delta[CH_n] / \Delta[CH_nC(=O)]$ $\sim -2/3$, and $\Delta[CH_n]/\Delta[CH_nONO_2] \sim 2/3$. This leads to a sorting of relative effectiveness of the groups in promoting miscibility, hydroxyl (-4) > acid (-2.5) > aldehyde (-2)> hydroperoxide (-1) > carbonyl (-0.66) > ether (-0.5)> nitrate (0.66), where the number in parentheses corresponds to the $\Delta[CH_n]/\Delta[n]$. According to this model the addition of nitrate groups is in the same direction as methylene groups; i.e., it reduces miscibility. This finding is consistent with CCN experiments on alkenes reacted with



Figure 3. Modeled κ values for homologous series of functionalized *n*-alkanes. Solid lines correspond to alkanes with 1–5 non-terminal hydroxyl groups. Orange dashed lines correspond to further functionalized dihydroxy alkanes as described in the legend. Colored carbon numbers (C₇, C₁₂, C₁₆, C₂₀, and C₂₄) correspond to the largest carbon number without miscibility-limited activation for the respective hydroxy alkanes series.

NO₃ radicals (Suda et al., 2014, their Supplement), and the known low miscibility of organic nitrates in water (Boschan et al., 1955). Furthermore, sorting of the different functional groups is qualitatively consistent with the sensitivity of κ to the addition of functional groups derived from CCN data (Table S5, Suda et al., 2014).

Treatment of OA evolution in the atmosphere

The computational speed of the model is relatively slow. The slow speed is due to the need to evaluate the entire range of mole fractions in order to determine the phase boundaries. Improvement in model execution speed is likely possible via algorithm optimization. Furthermore, parallel execution of the code is possible. With a regular workstation it is feasible to perform offline computation of $\sim 10^6 \kappa$'s for a large set of compounds produced by the Generator of Explicit Chemistry and Kinetics of Organics in the Atmosphere (GECKO-A) or similar models. Once pure component κ 's are predicted, the evolution of the overall OA κ in mixed particles can be calculated quickly using the linear mixing rule (Petters and Kreidenweis, 2007), subject to the limitations of this approach discussed in Sect. 2. One additional limitation is the need for algorithms that automatically map the computer-generated simplified molecularinput line-entry system (SMILES) structures (e.g., Table 3 in Lee-Taylor et al., 2015) to UNIFAC groups. Several of these structures are bridged and even manual mapping of those structures to UNIFAC groupings will necessitate definition of new groups with unknown volume, surface, and interaction parameters. Separate studies are needed to establish the minimal number of new groups that would be needed to obtain optimal coverage for the set of compounds of interest.

4 Summary and conclusions

This paper describes how functional group contribution methods can be used to estimate the CCN activity of pure organic compounds. Group interaction parameters were taken from a mix of sources and used without further tuning. Model fidelity was evaluated against a database of published CCN data. Weakly functionalized alkanes are correctly classified as effectively CCN inactive (defined as $\kappa < 0.001$). Highly functionalized and water-soluble molecules are predicted to activate in accordance with the estimated molar volume and generally predictions agree with observations within a factor of 2. Liquid-liquid phase separation is predicted to occur for compounds with few functional groups and phase separation is predicted to control κ . The model adequately reproduces the observation that hydroxyl groups strongly promote CCN activity while nitrate groups inhibit CCN activity. A few outliers in the model evaluation may be explained by the combination of CCN measurement uncertainty, compound purity, uncertainty in dry particle phase state, and insufficiently tuned group interaction parameters. However, more systematic data on weakly functionalized compounds, including repeat studies, are needed before a retuning of parameters is justified. The model makes new predictions about the relative effectiveness of the groups in promoting miscibility. Most notably, it predicts that hydroperoxides have much less of an effect than hydroxyl, which is slightly surprising since one would expect the hydrogen bonding to be similar. The model state space can serve as a rough guide to define test conditions to quantify via experiment the effectiveness of adding one or more functional groups to a carbon backbone.

Although this work is limited to a few functional groups, the presented framework is general since interaction parameters are available for a wide range of groups. For atmospheric purposes, amines, olefins, and aromatic compounds are the most relevant groups that need to be added. Few, if any, systematic CCN data for these groups are available. However, the success of the current model to estimate κ without the need to tune parameters could be taken as an indication that first-order predictions can be obtained until such data become available.

5 Code availability

Source code and example scripts demonstrating model initialization for the compounds presented in this study are available as Supplement to this manuscript.

The Supplement related to this article is available online at doi:10.5194/gmd-9-111-2016-supplement.

Acknowledgements. This work was funded by the Department of Energy, Office of Biological and Environmental Sciences under grant DE-SC0010470.

Edited by: A. Archibald

References

- Amundson, N. R., Caboussat, A., He, J. W., and Seinfeld, J. H.: An optimization problem related to the modeling of atmospheric organic aerosols, C. R. Acad. Sci. Paris Ser. I, 340, 765–768, doi:10.1016/j.crma.2005.04.018, 2005.
- Amundson, N. R., Caboussat, A., He, J. W., Martynenko, A. V., Landry, C., Tong, C., and Seinfeld, J. H.: A new atmospheric aerosol phase equilibrium model (UHAERO): organic systems, Atmos. Chem. Phys., 7, 4675–4698, doi:10.5194/acp-7-4675-2007, 2007.
- Andreae, M. O. and Rosenfeld, D.: Aerosol-cloudprecipitation interactions, Part 1. The nature and sources

of cloud-active aerosols, Earth-Sci. Rev., 89, 13–41, doi:10.1016/j.earscirev.2008.03.001, 2008.

- Barley, M. H., Topping, D. O., and McFiggans, G.: The critical assessment of liquid density estimation methods for multifunctional organic compounds, J. Phys. Chem. A, 117, 3428–3441, doi:10.1021/jp304547r, 2013.
- Bilde, M. and Svenningsson, B.: CCN activation of slightly soluble organics: The importance of small amounts of inorganic salt and particle phase, Tellus, 56B, 128–134, doi:10.1111/j.1600-0889.2004.00090.x, 2004
- Bondi, A.: van der Waals Volumes and Radii, J. Phys. Chem., 68, 441–451, doi:10.1021/j100785a001, 1964.
- Boschan, R., Merrow, R. T., and van Dolah, R. W.: The chemistry of nitrate esters, Chem. Rev., 55, 485–510, doi:10.1021/cr50003a001, 1955.
- Broekhuizen, K., Kumar, P. P., and Abbatt, J. P. D.: Partially soluble organics as cloud condensation nuclei: Role of trace soluble and surface active species, Geophys. Res. Lett., 31, L01107, doi:10.1029/2003GL018203, 2004.
- Cappa, C. D., Che, D. L., Kessler, S. H., Kroll, J. H., and Wilson, K. R.: Variations in organic aerosol optical and hygroscopic properties upon heterogeneous OH oxidation, J. Geophys. Res., 116, D15204, doi:10.1029/2011JD015918, 2011.
- Chang, R. Y.-W., Slowik, J. G., Shantz, N. C., Vlasenko, A., Liggio, J., Sjostedt, S. J., Leaitch, W. R., and Abbatt, J. P. D.: The hygroscopicity parameter (κ) of ambient organic aerosol at a field site subject to biogenic and anthropogenic influences: relationship to degree of aerosol oxidation, Atmos. Chem. Phys., 10, 5047–5064, doi:10.5194/acp-10-5047-2010, 2010.
- Choi, M. Y. and Chan, C. K.: Continuous measurements of the water activities of aqueous droplets of water-soluble organic compounds, J. Phys. Chem. A., 106, 4566–4572, doi:10.1021/jp0138750, 2002.
- Christensen, S. I. and Petters, M. D.: The role of temperature in cloud droplet activation, J. Phys. Chem. A., 116, 9706–9717, doi:10.1021/jp3064454, 2012.
- Clegg, S. L. and Seinfeld, J. H.: Thermodynamic models of aqueous solutions containing inorganic electrolytes and dicarboxylic acids at 298.15 K. I., The acids as non-dissociating components, J. Phys. Chem. A, 110, 5692–5717, doi:10.1021/jp056150j, 2006.
- Compernolle, S., Ceulemans, K., and Müller, J.-F.: Influence of non-ideality on condensation to aerosol, Atmos. Chem. Phys., 9, 1325–1337, doi:10.5194/acp-9-1325-2009, 2009.
- Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmpadimos, I., Prevot, A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez, J. L., Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and Baltensperger, U.: Relating hygroscopicity and composition of organic aerosol particulate matter, Atmos. Chem. Phys., 11, 1155– 1165, doi:10.5194/acp-11-1155-2011, 2011.
- Eubank, P. T., Elhassan, A. E., Barrufet, M. A., and Whiting, W. B.: Area method for prediction of fluid-phase equilibria, Ind. Eng. Chem. Res., 31, 942–949, doi:10.1021/ie00003a041,1992.
- Fredenslund, A., Jones, R. L., and Prausnitz, J. M.: Group-contribution estimation of activity-coefficients in nonideal liquid-mixtures, AIChE J., 21, 1086–1099, doi:10.1002/aic.690210607, 1975.

- George, I. J. and Abbatt, J. P. D.: Heterogeneous oxidation of atmospheric aerosol particles by gas-phase radicals, Nature Chem., 2, 713–722, doi:10.1038/nchem.806, 2010.
- Girolami G. S.: A Simple "Back of the Envelope" Method for Estimating the Densities and Molecular Volume of Liquids and Volumes, J. Chem. Education, 71, 962–964, doi:10.1021/ed071p962, 1994.
- Goldstein, A. H. and Galbally, I. E.: Known and Unexplored Organic Constituents in the Earth's Atmosphere, Environ. Sci. Technol., 41, 1514–1521, doi:10.1021/es072476p, 2007.
- Hansen, H. K., Rasmussen, P., Fredenslund, A., Schiller, M., and Gmehling, J.: Vapor-liquid equilibria by UNIFAC group contribution. 5. Revision and extension, Ind. Eng. Chem. Res., 30, 2352–2355, doi:10.1021/ie00058a017, 1991.
- Heald, C. L., Henze, D. K., Horowitz, L. W., Feddema, J., Lamarque, J. F., Guenther, A., Hess, P. G., Vitt, F., Seinfeld, J. H., Goldstein, A. H., and Fung, I.: Predicted change in global secondary organic aerosol concentrations in response to future climate, emissions, and land use change, J. Geophys. Res., 113, D05211, doi:10.1029/2007jd009092, 2008.
- Hings, S. S., Wrobel, W. C., Cross, E. S., Worsnop, D. R., Davidovits, P., and Onasch, T. B.: CCN activation experiments with adipic acid: effect of particle phase and adipic acid coatings on soluble and insoluble particles, Atmos. Chem. Phys., 8, 3735– 3748, doi:10.5194/acp-8-3735-2008, 2008.
- Hori, M., Ohta, S., Murao, N., and Yamagata, S.: Activation capability of water soluble organic substances as CCN, J. Aerosol. Sci., 34, 419–448, doi:10.1016/S0021-8502(02)00190-8, 2003.
- Huff Hartz, K. E., Tischuk, J. E., Chan, M. N., Chan, C. K., Donahue, N. M., and Pandis, S. N.: Cloud condensation nuclei activation of limited solubility organic aerosol, Atmos. Environ., 40, 605–617, doi:10.1016/j.atmosenv.2005.09.076, 2006.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J.D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina. K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, Science, 326, 1525-1529, doi:10.1126/science.1180353, 2009.
- Joback, K. G. and Reid, R. C.: Estimation of pure-component properties from group-contributions, Chem. Eng. Commun., 57, 233– 243, doi:10.1080/00986448708960487, 1987.
- Kreidenweis, S. M., Petters, M. D., and DeMott, P. J.: Deliquescence-controlled activation of organic aerosols, Geophys. Res. Lett., 33, L06801, doi:10.1029/2005GL024863, 2006.
- Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of low-volatility organics in the atmosphere, Atmos. Environ., 42, 3593–3624, doi:10.1016/j.atmosenv.2008.01.003, 2008.

- Kuwata, M., Zorn, S. R., and Martin, S. T.: Using elemental ratios to predict the density of 167 organic material composed of carbon, hydrogen, and oxygen, Environ. Sci. Technol., 46, 787–794, doi:10.1021/es202525q, 2012.
- Kuwata, M., Shao, W., Lebouteiller, R., and Martin, S. T.: Classifying organic materials by oxygen-to-carbon elemental ratio to predict the activation regime of Cloud Condensation Nuclei (CCN), Atmos. Chem. Phys., 13, 5309–5324, doi:10.5194/acp-13-5309-2013, 2013.
- Lambe, A. T., Onasch, T. B., Massoli, P., Croasdale, D. R., Wright, J. P., Ahern, A. T., Williams, L. R., Worsnop, D. R., Brune, W. H., and Davidovits, P.: Laboratory studies of the chemical composition and cloud condensation nuclei (CCN) activity of secondary organic aerosol (SOA) and oxidized primary organic aerosol (OPOA), Atmos. Chem. Phys., 11, 8913–8928, doi:10.5194/acp-11-8913-2011, 2011.
- Lee-Taylor, J., Hodzic, A., Madronich, S., Aumont, B., Camredon, M., and Valorso, R.: Multiday production of condensing organic aerosol mass in urban and forest outflow, Atmos. Chem. Phys., 15, 595–615, doi:10.5194/acp-15-595-2015, 2015.
- Li, Z., Williams, A. L., and Rood, M. J.: Influence of soluble surfactant properties on the activation of aerosol particles containing inorganic solute, J. Atmos. Sci., 55, 1859–1866, 1998.
- Liu, X. and Wang, J.: How important is organic aerosol hygroscopicity to aerosol indirect forcing, Environ. Res. Lett., 5, 044010, doi:10.1088/1748-9326/5/4/044010, 2010.
- Marcolli, C. and Peter, Th.: Water activity in polyol/water systems: new UNIFAC parameterization, Atmos. Chem. Phys., 5, 1545– 1555, doi:10.5194/acp-5-1545-2005, 2005.
- Marcolli, C., Luo, B. P., and Peter, T.: Mixing of the organic aerosol fractions: Liquids as the thermodynamically stable phases, J. Phys. Chem. A, 108, 2216–2224, doi:10.1021/jp036080l, 2004.
- Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkila, J., Canagaratna, M. R., Brune, W. H., Onasch, T. B., Jayne, J. T., Petaja, T., Kulmala, M., Laaksonen, A., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Relationship between aerosol oxidation level and hygroscopic properties of laboratory generated secondary organic aerosol (SOA) particles, Geophys. Res. Lett., 37, L24801, doi:10.1029/2010GL045258, 2010.
- Mei, F., Setyan, A., Zhang, Q., and Wang, J.: CCN activity of organic aerosols observed downwind of urban emissions during CARES, Atmos. Chem. Phys., 13, 12155–12169, doi:10.5194/acp-13-12155-2013, 2013.
- Ming, Y. and Russell, L. M.: Predicted hygroscopic growth of sea salt aerosol, J. Geophys. Res., 106, 28259–28274, doi:10.1029/2001JD000454, 2001.
- Ming, Y. and Russell, L. M.: Thermodynamic equilibrium of organic–electrolyte mixtures in aerosol particles, AIChE J., 48, 1331–1348, doi:10.1002/aic.690480619, 2002.
- Mochida, M. and Kawamura, K.: Hygroscopic properties of levoglucosan and related organic compounds characteristic to biomass burning aerosol particles, J. Geophys. Res., 109, D21202, doi:10.1029/2004JD004962, 2004.
- Moore, R. H. and Raymond, T. M.: HTDMA analysis of multicomponent dicarboxylic acid aerosols with comparison to UNIFAC and ZSR, J. Geophys. Res.-Atmos., 113, D04206, doi:10.1029/2007jd008660, 2008.
- Nguyen, T. K. V., Petters, M. D., Suda, S. R., Guo, H., Weber, R. J., and Carlton, A. G.: Trends in particle-phase liquid water during

M. D. Petters et al.: Prediction of cloud condensation nuclei activity for organic compounds

the Southern Oxidant and Aerosol Study, Atmos. Chem. Phys., 14, 10911–10930, doi:10.5194/acp-14-10911-2014, 2014.

- Peng, C., Chan, M. N., and Chan, C. K.: The hygroscopic properties of dicarboxylic and multifunctional acids: Measurements and UNIFAC predictions, Environ. Sci. Technol., 35, 4495–4501, doi:10.1021/es0107531, 2001.
- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961–1971, doi:10.5194/acp-7-1961-2007, 2007.
- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud condensation nucleus activity – Part 2: Including solubility, Atmos. Chem. Phys., 8, 6273–6279, doi:10.5194/acp-8-6273-2008, 2008.
- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud condensation nucleus activity – Part 3: Including surfactant partitioning, Atmos. Chem. Phys., 13, 1081–1091, doi:10.5194/acp-13-1081-2013, 2013.
- Petters, M. D., Kreidenweis, S. M., Snider, J. R., Koehler, K. A., Wang, Q., and Prenni, A. J.: Cloud droplet activation of polymerized organic aerosol, Tellus B, 58, 195–205, doi:10.1111/j.1600-0889.2006.00181.x, 2006.
- Petters, M. D., Wex, H., Carrico, C. M., Hallbauer, E., Massling, A., McMeeking, G. R., Poulain, L., Wu, Z., Kreidenweis, S. M., and Stratmann, F.: Towards closing the gap between hygroscopic growth and activation for secondary organic aerosol – Part 2: Theoretical approaches, Atmos. Chem. Phys., 9, 3999– 4009, doi:10.5194/acp-9-3999-2009, 2009a.
- Petters, M. D., Kreidenweis, S. M., Prenni, A. J., Sullivan, R. C., Carrico, C. M., Koehler, K. A., and Ziemann, P. J.: Role of molecular size in cloud droplet activation, Geophys. Res Lett., 36, L22801, doi:10.1029/2009GL040131, 2009b.
- Poulain, L., Wu, Z., Petters, M. D., Wex, H., Hallbauer, E., Wehner, B., Massling, A., Kreidenweis, S. M., and Stratmann, F.: Towards closing the gap between hygroscopic growth and CCN activation for secondary organic aerosols – Part 3: Influence of the chemical composition on the hygroscopic properties and volatile fractions of aerosols, Atmos. Chem. Phys., 10, 3775– 3785, doi:10.5194/acp-10-3775-2010, 2010.
- Prausnitz, J. M., Lichtenthaler, R. N., and Azevedo, E. G.: Molecular thermodynamics of fluid-phase equilibria, Prentice Hall PTR, Upper Saddle River, NJ, USA, 1999.
- Prisle, N. L., Dal Maso, M., and Kokkola, H.: A simple representation of surface active organic aerosol in cloud droplet formation, Atmos. Chem. Phys., 11, 4073–4083, doi:10.5194/acp-11-4073-2011, 2011.
- Raatikainen, T. and Laaksonen, A.: Application of several activity coefficient models to water-organic-electrolyte aerosols of atmospheric interest, Atmos. Chem. Phys., 5, 2475–2495, doi:10.5194/acp-5-2475-2005, 2005.
- Raatikainen, T. and Laaksonen, A.: A simplified treatment of surfactant effects on cloud drop activation, Geosci. Model Dev., 4, 107–116, doi:10.5194/gmd-4-107-2011, 2011.
- Rafati, A. A., Bagheri, A., Khanchi, A. R., Ghasemian, E., and Najafi, M.: Application of the UNIFAC model for prediction of surface tension and thickness of the surface layer in the binary mixtures, J. Colloid Interf. Sci., 355, 252–258, doi:10.1016/j.jcis.2010.12.003, 2011.

- Raymond, T. M. and Pandis, S. N.: Cloud activation of singlecomponent organic aerosol particles, J. Geophys. Res., 107, 4787, doi:4710.1029/2002JD002159, 2002.
- Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, M., Sellier, M., Murray, B. J., Shilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of α -pinene secondary organic material and implications for particle growth and reactivity, P. Natl. Acad. Sci., 110, 8014–8019, doi:10.1073/pnas.1219548110, 2013.
- Rickards, A. M. J., Miles, R. E. H., Davies, J. F., Marshall, F. H., and Reid, J. P.: Measurements of the sensitivity of aerosol hygroscopicity and the k parameter to the O / C Ratio, J. Phys. Chem. A, 117, 14120–14131, doi:10.1021/jp407991n, 2013.
- Rissman, T. A., Varutbangkul, V., Surratt, J. D., Topping, D. O., McFiggans, G., Flagan, R. C., and Seinfeld, J. H.: Cloud condensation nucleus (CCN) behavior of organic aerosol particles generated by atomization of water and methanol solutions, Atmos. Chem. Phys., 7, 2949–2971, doi:10.5194/acp-7-2949-2007, 2007.
- Rood, M. J. and Williams, A. L.: Reply, J. Atmos. Sci., 58, 1468– 1473, 2001.
- Rood, M. J., Shaw, M. A., Larson, T. V., and Covert, D. S.: Ubiquitous nature of ambient metastable aerosol, Nature, 337, 537–539, 1989.
- Sastri, S. R. S. and Rao, K. K.: A new group contribution method for predicting viscosity of organic liquids, Chem. Eng. J., 50, 9–25, doi:10.1016/0300-9467(92)80002-R, 1992.
- Sastri, S. R. S. and Rao, K. K.: A simple method to predict surface tension of organic liquids, Chem. Eng. J. Bioch. Eng., 59, 181– 186, doi:10.1016/0923-0467(94)02946-6, 1995.
- Saxena, P. and Hildemann, L. M.: Water absorption by organics: Survey of laboratory evidence and evaluation of UNIFAC for estimating water activity, Environ. Sci. Technol., 31, 3318–3324, doi:10.1021/es9703638, 1997.
- Seinfeld, J. and Pandis, S.: Atmospheric chemistry and physics: from air pollution to climate change, Wiley, Inc., New Jersey, USA, 450 pp., 2006.
- Shilling, J. E., King, S. M., Mochida, M., and Martin, S. T.: Mass spectral evidence that small changes 574 in composition caused by oxidative aging processes alter aerosol CCN properties, J. Phys. Chem., 111, 3358–3368, doi:10.1021/jp068822r, 2007.
- Shiraiwa, M., Ammann, M., Koop, T., and Pöschl, U.: Gas uptake and chemical aging of semisolid organic aerosol particles, P. Natl. Acad. Sci., 108, 11003–11008, doi:10.1073/pnas.1103045108, 2011.
- Shulman, M. L., Jacobson, M. C., Charlson, R. J., Synovec, R. E., and Young, T. E.: Dissolution behavior and surface tension effects of organic compounds in nucleating cloud droplets, Geophys. Res. Lett., 23, 277–280, doi:10.1029/95GL03810, 1996.
- Sorjamaa, R., Svenningsson, B., Raatikainen, T., Henning, S., Bilde, M., and Laaksonen, A.: The role of surfactants in Köhler theory reconsidered, Atmos. Chem. Phys., 4, 2107–2117, doi:10.5194/acp-4-2107-2004, 2004.
- Sprow, F. B. and Prausnitz, J. M.: Surface tensions of simple liquids, T. Faraday Soc., 62, 1097–1104, doi:10.1039/TF9666201105, 1966.
- Suda, S. R., Petters, M. D., Yeh, K., Strollo, C., Matsunaga, A., Faulhaber, A., Ziemann, P. J., Prenni, A. J., Carrico, C. M., Sullivan, R. C., and Kreidenweis, S. M.: Influence of functional

groups on organic aerosol cloud condensation nucleus activity, Environ. Sci. Technol., 48, 10182–10190, 2014.

- Topping, D., Barley, M., and McFiggans G.: Including phase separation in a unified model to calculate partitioning of vapours to mixed inorganic–organic aerosol particles, Faraday Discuss., 165, 273–288, doi:10.1039/C3FD00047H, 2013.
- Topping, D. O., McFiggans, G. B., and Coe, H.: A curved multicomponent aerosol hygroscopicity model framework: Part 2 – Including organic compounds, Atmos. Chem. Phys., 5, 1223–1242, doi:10.5194/acp-5-1223-2005, 2005.
- Tuckermann, R.: Surface tension of aqueous solutions of watersoluble organic and inorganic compounds, Atmos. Environ., 41, 6265–6275, 2007.
- Tuckermann, R. and Cammenga, H. K.: The surface tension of aqueous solutions of some atmospheric water-soluble organic compounds, Atmos. Environ., 38, 6135–6138, 2004.
- Vaden, T. D., Imre, D., Beránek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation kinetics and phase of laboratory and ambient secondary organic aerosol, P. Natl. Acad. Sci., 108, 2190–2195, doi:10.1073/pnas.1013391108, 2011.

- Ziemann, P. J. and Atkinson, R: Kinetics, products, and mechanisms of secondary organic aerosol formation, Chem. Soc. Rev., 41, 6582–6605, doi:10.1039/C2CS35122F, 2012.
- Zobrist, B., Soonsin, V., Luo, B. P., Krieger, U. K., Marcolli, C., Peter, T., and Koop, T.: Ultra-slow water diffusion in aqueous sucrose glasses, Phys. Chem. Chem. Phys., 13, 3514–3526, doi:10.1039/c0cp01273d, 2011.
- Zuend, A. and Seinfeld, J. H.: Modeling the gas-particle partitioning of secondary organic aerosol: the importance of liquidliquid phase separation, Atmos. Chem. Phys., 12, 3857–3882, doi:10.5194/acp-12-3857-2012, 2012.
- Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organic-inorganic aerosols to predict activity coefficients, Atmos. Chem. Phys., 8, 4559–4593, doi:10.5194/acp-8-4559-2008, 2008.
- Zuend, A., Marcolli, C., Peter, T., and Seinfeld, J. H.: Computation of liquid-liquid equilibria and phase stabilities: implications for RH-dependent gas/particle partitioning of organic-inorganic aerosols, Atmos. Chem. Phys., 10, 7795–7820, doi:10.5194/acp-10-7795-2010, 2010.

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